

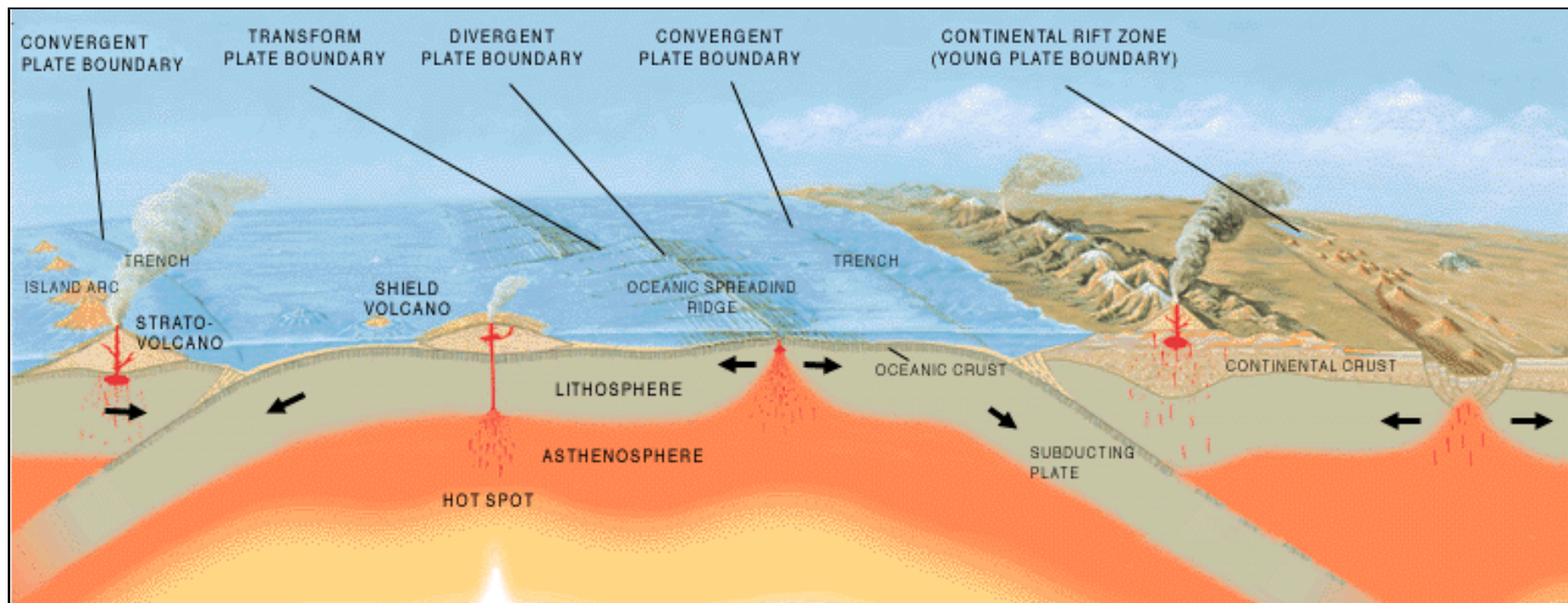
# Vertical Movements

## Geology 1

G. Bertotti

Lateral deformations can change the **thickness of the lithospheric column**.

This has major implications **vertical movements** (**subsidence and uplift**) of the Earth's surface leading to sedimentary basins and mountains



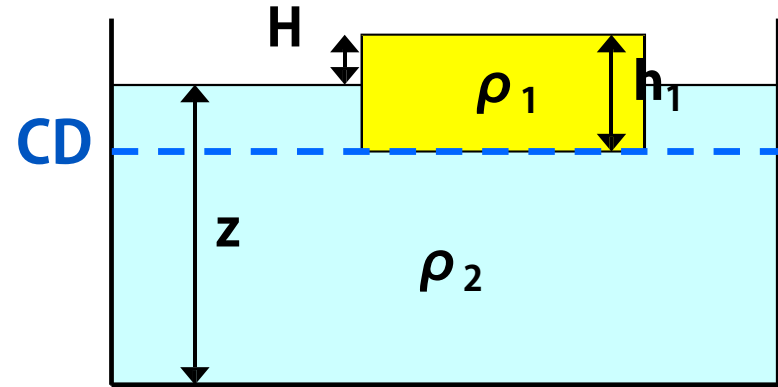
**Isostasy:**

linking weight to **topography**

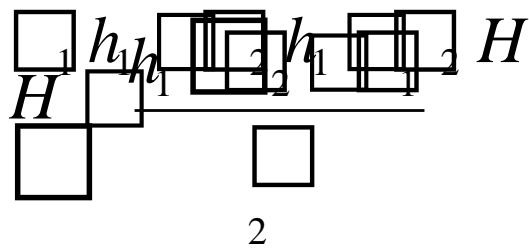
Linking weight changes to **vertical movements**

Isostasy is **Archimedes' law:**

The weight of the floating body must be equal to that of the displaced body



$$\rho_1 g h_1 = \rho_2 g (h_1 - H)$$

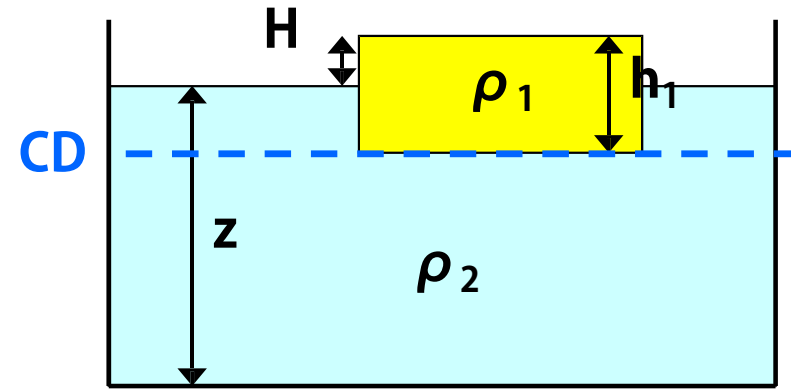


An equivalent expression:  
the weight of the column above a **compensation depth** must be the same  
CD=the depth under which ..nothing happens

$$(h_1 - H) \rho_2 = \rho_1 h_1$$

H is the **topography**, the number we are interested in

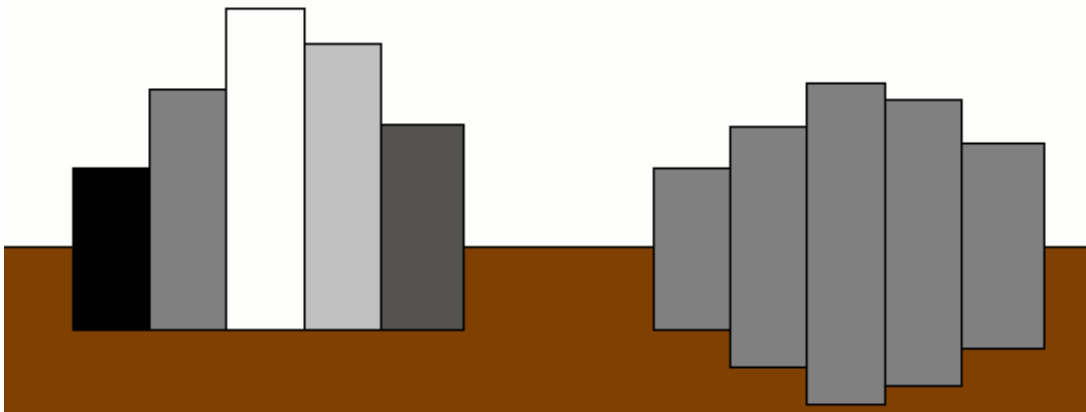
Lateral changes in topography controlled by changes in the weight of the column overlying the CD



$$H = \frac{h_1 (\rho_1 + \rho_2)}{\rho_2}$$

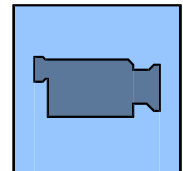
The lower the density of the column, the higher the mountains

The higher the thickness of the column, the higher will be the topography



One can change densities (left), thicknesses (right) or both

<http://www.geo.cornell.edu/hawaii/220/PRI/isostasy.html>

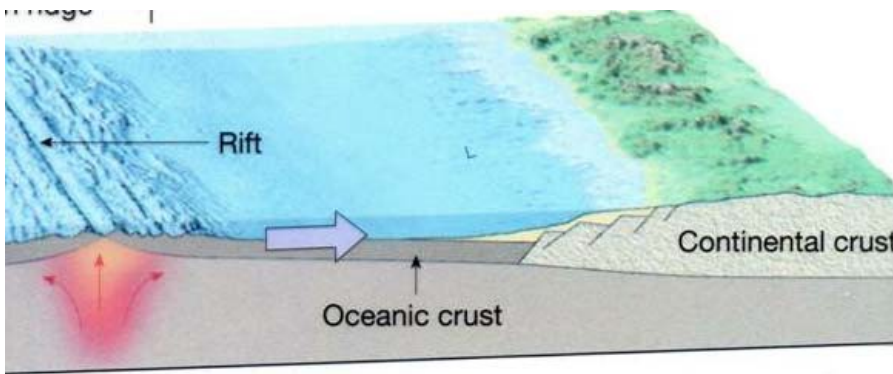
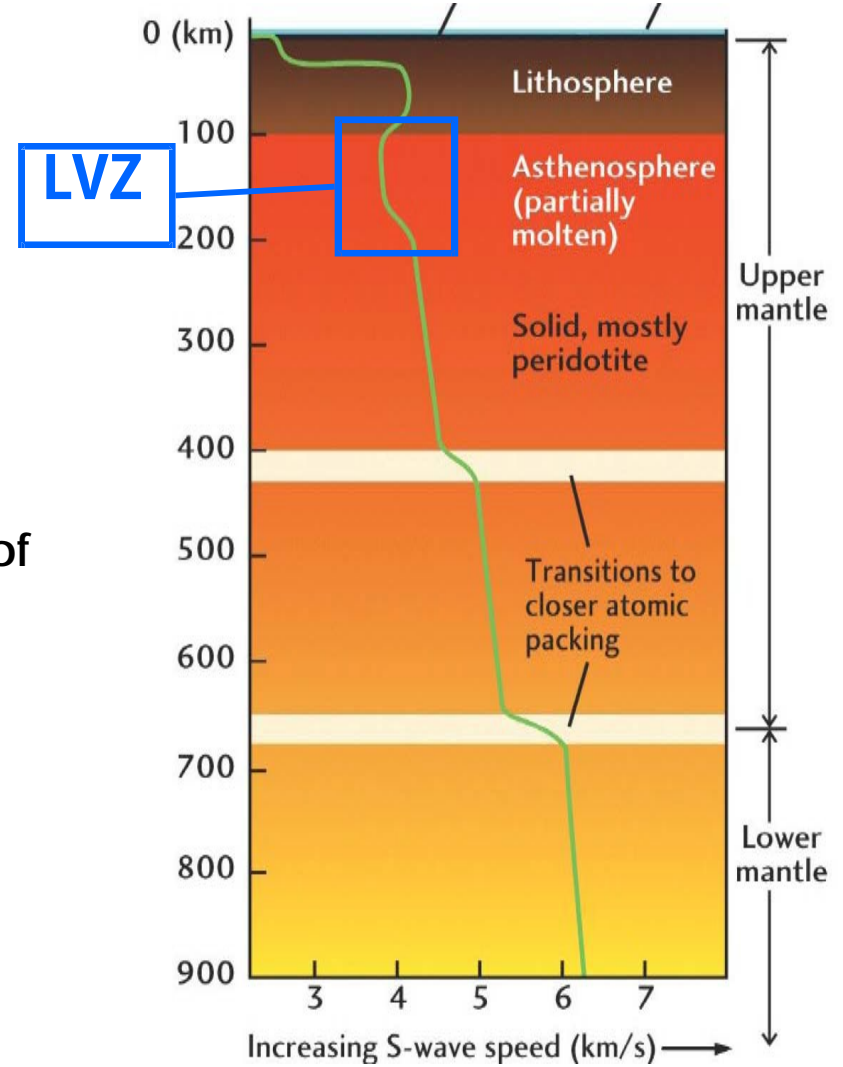


What about the Earth:  
 Where is the **compensation level**?  
 which is the **floating block**?

The best candidate for the compensation level is the **LVZ** with the **lithosphere** as floating block

If (and this seems to be true) it floats, the density of the lithosphere must be less than that of the asthenosphere

$$\rho_{\text{lith}} < \rho_{\text{asth.}}$$

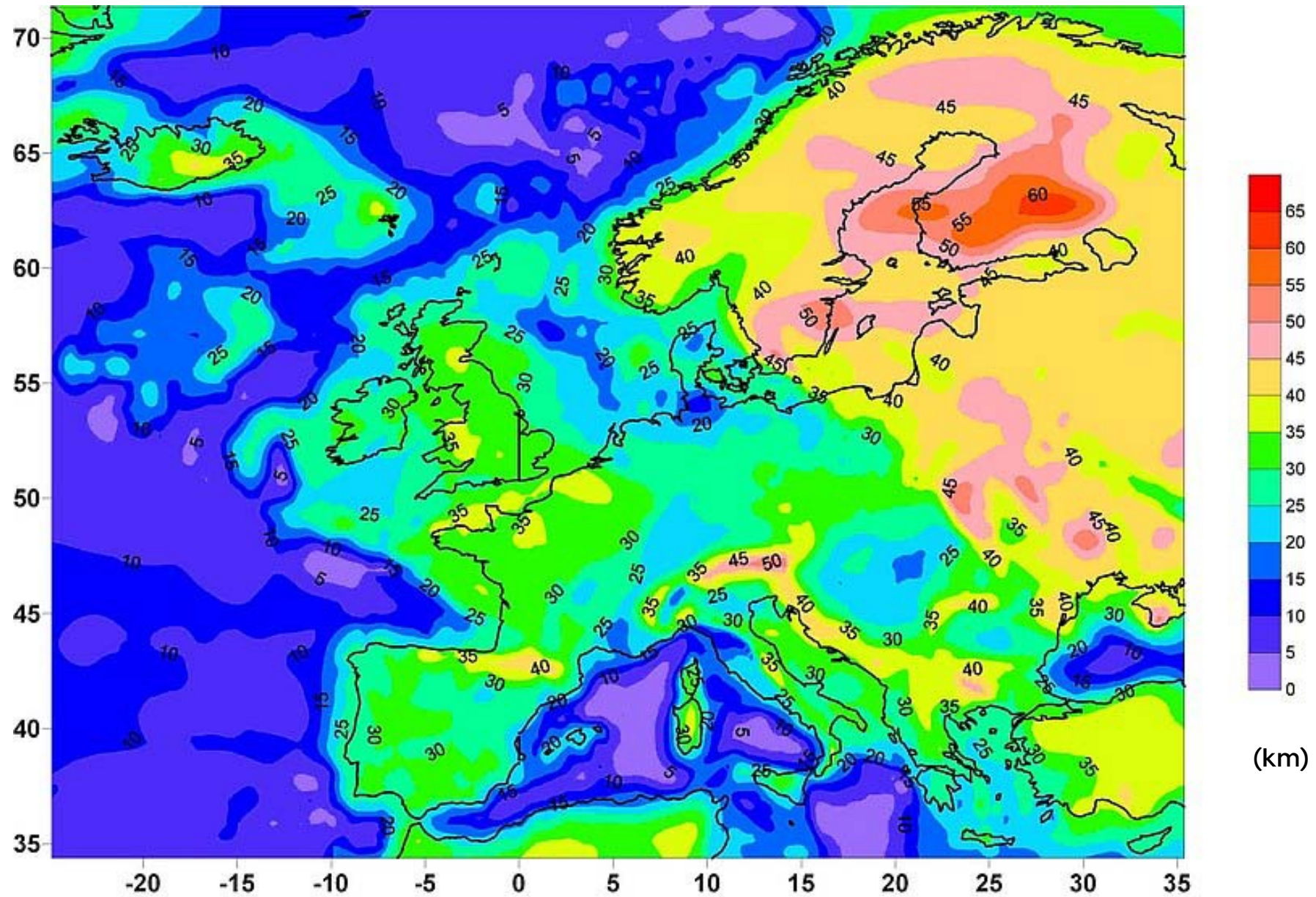


Remember that the lithosphere is a two components system, lith. mantle and (**oceanic-continental**) crust





## The Moho map of Europe







# Changing weights (= vertical movements) in the Earth

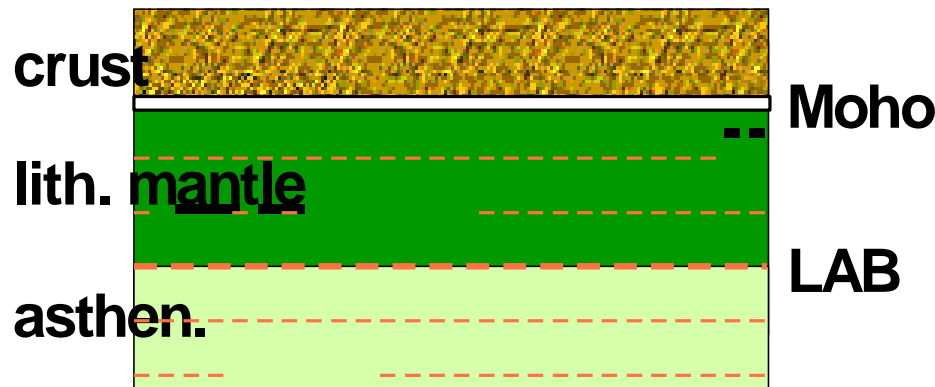
$$\rho_{\text{lith}} < \rho_{\text{asth.}}$$

This is what we had. Life seems easy

If the lithosphere becomes **lighter** topography **increases**

If lithosphere becomes **heavier** topography **decreases**

*Geologically not really useful, the lithosphere is a complicated “block”*



$$\rho_{\text{crust}} = 2.6 \text{ g} \cdot \text{cm}^{-3}$$

$$\rho_{\text{lith.mant}} = 3.6 \text{ g} \cdot \text{cm}^{-3}$$

$$\rho_{\text{asthen}} = 3.0 \text{ g} \cdot \text{cm}^{-3}$$

$$\rho_{\text{crust}} < \rho_{\text{asth.}} < \rho_{\text{lith mantle}}$$

It seems that crust and lithospheric mantle play a different game!

We need to look in more detail processes controlling their thickness changes and add them together (quantitative geology!)

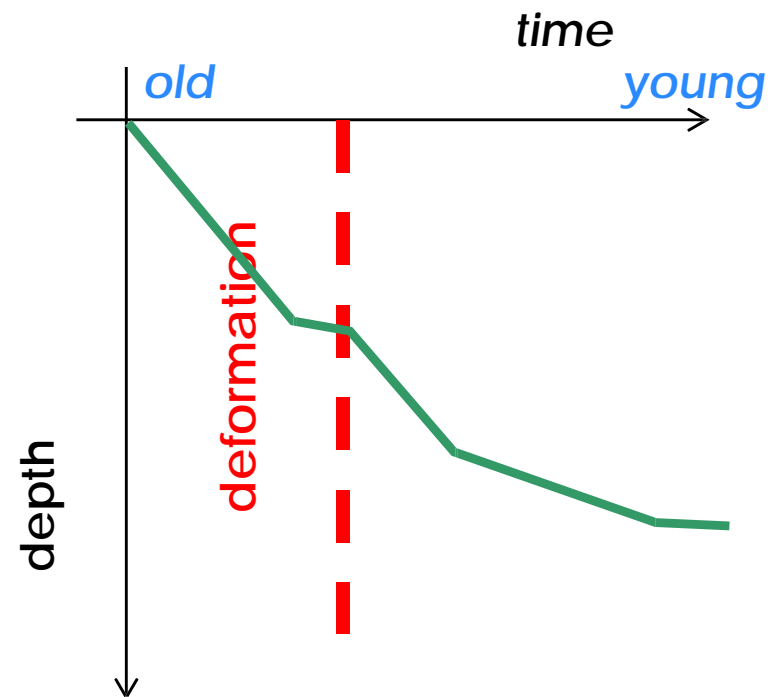
# We make a **conceptual** experiment (= simpler than nature)

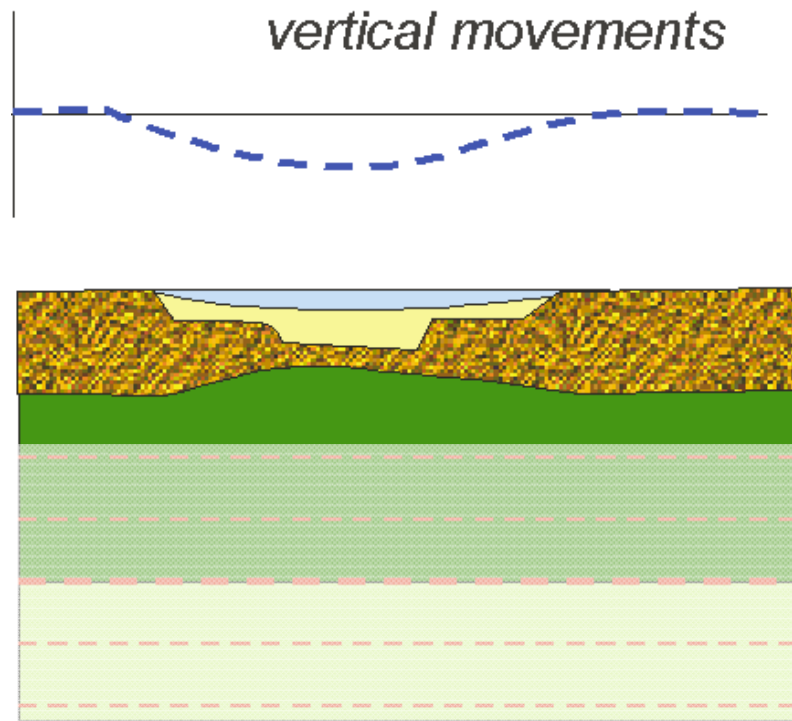
We investigate the vertical movements produced by thinning/thickening of the crust and lithospheric mantle **separately**

We do this by predicting **vertical movements**

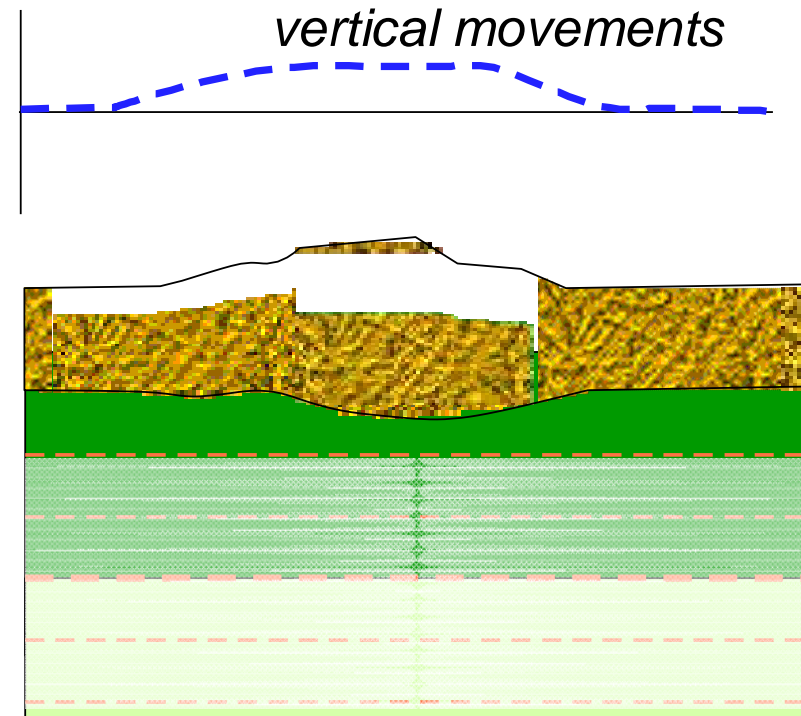
- 1) during thickness changes
- 2) **after** thickness changes

A useful tool to describe vertical movements through time is a **subsidence curve**



Effects of changes in **crustal thickness**

thinning

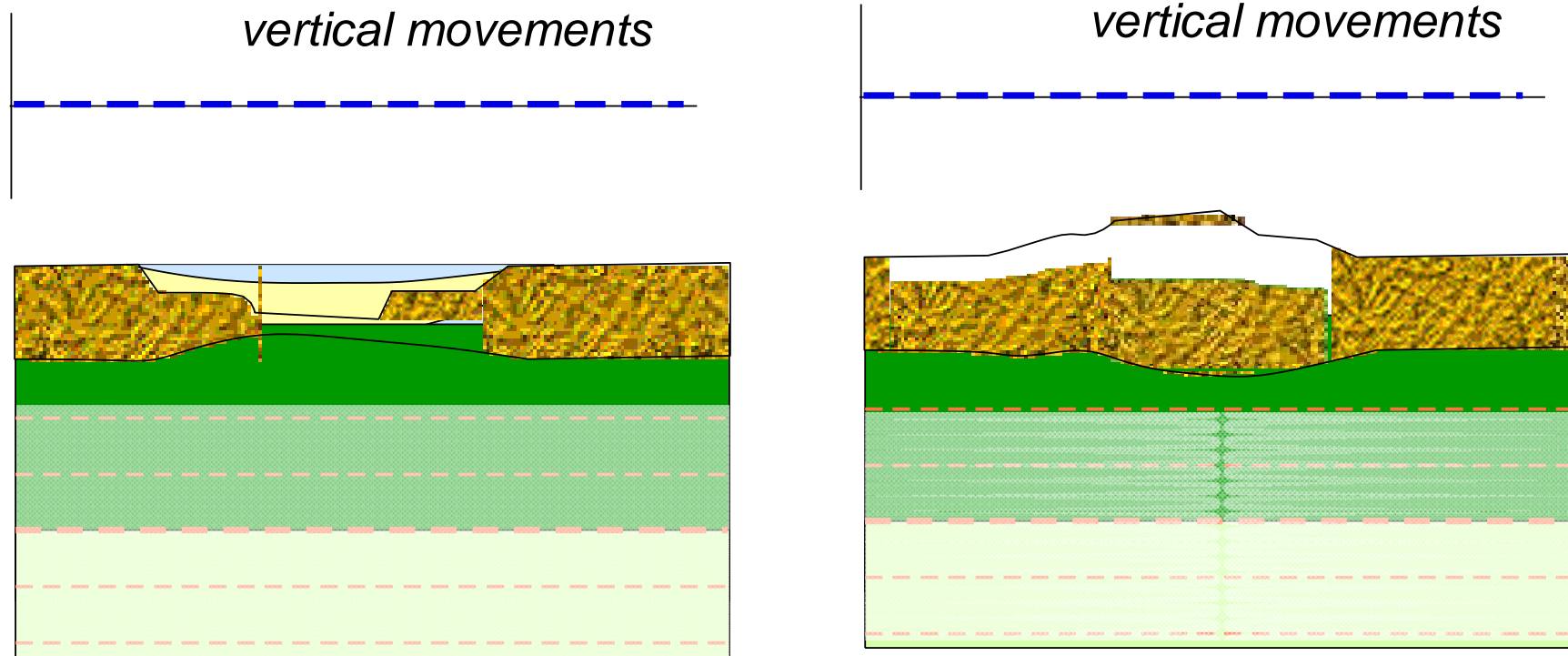


thickening

Easy:

- *crustal thinning causes downward movement (=subsidence)*
- *crustal thickening causes uplift*

After the end of crustal thickness changes....

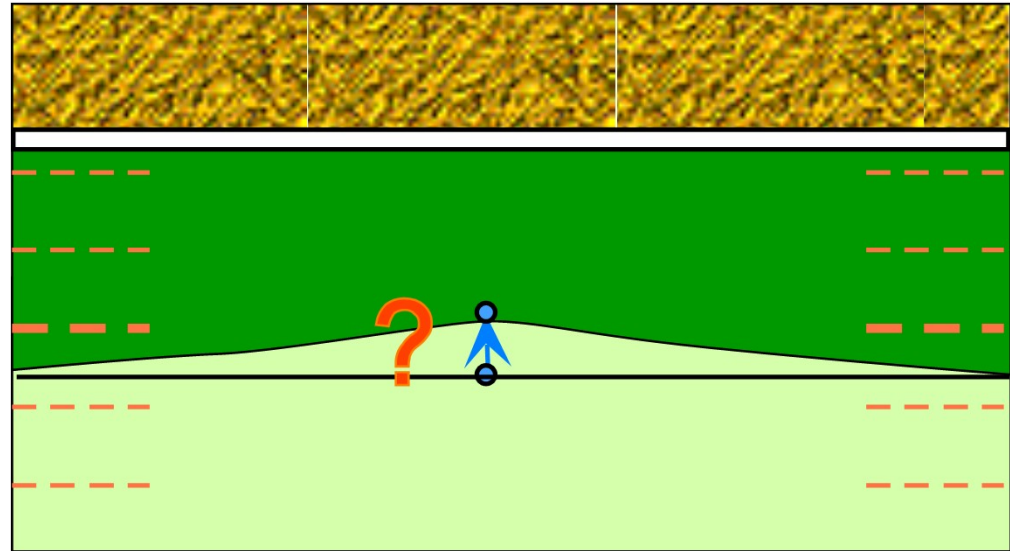


Nothing happens, no **crust-driven** vertical movements



## The lithospheric mantle

More complicated as the base of the lithosphere is a thermal boundary



In terms of mass movements

- During “extension/thinning”, mantle rocks will tend to move **upward**
- during “contraction/thickening”, mantle rocks will tend to move **downward**

But

If the mantle changes in thickness or not depends on the **competition** between **advection** (movement of mass) and **conduction** (movement of **heat**)

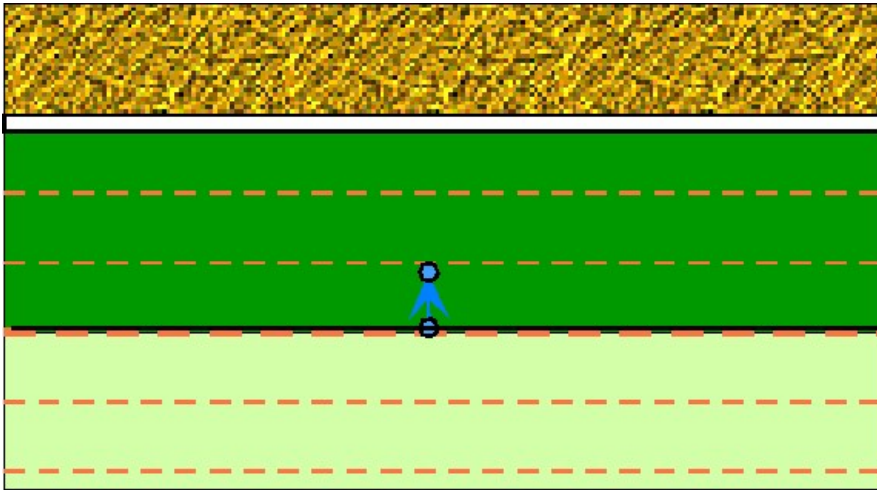
The competition  
between conduction  
and advection:

You better sail fast!



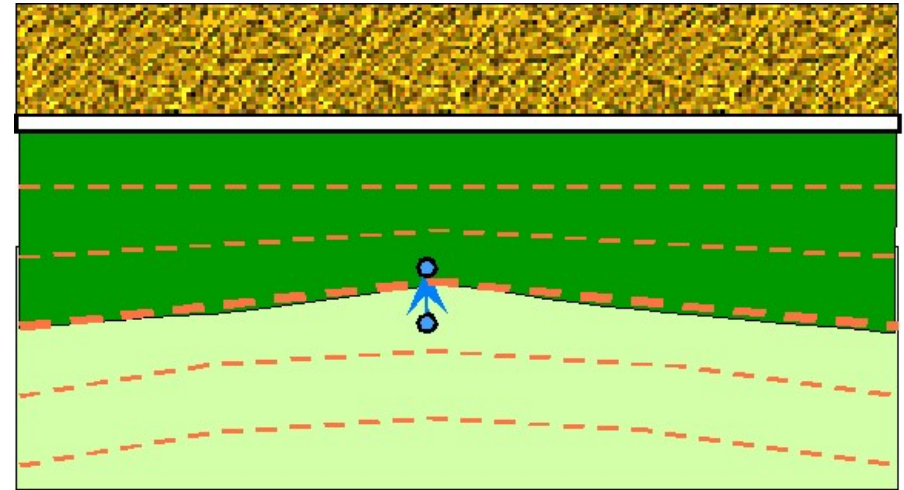
In the case of extension

If **conduction** prevails isotherms remain flat and the lithosphere does not thin



when thinning is **slow**  
 conduction  $\gg$  advection  $\blacktriangleright$  little  
 thermal anomaly  $\blacktriangleright$  **little**  
**lithospheric thinning**

If **advection** prevails isotherms are deflected and the lithosphere becomes thinner



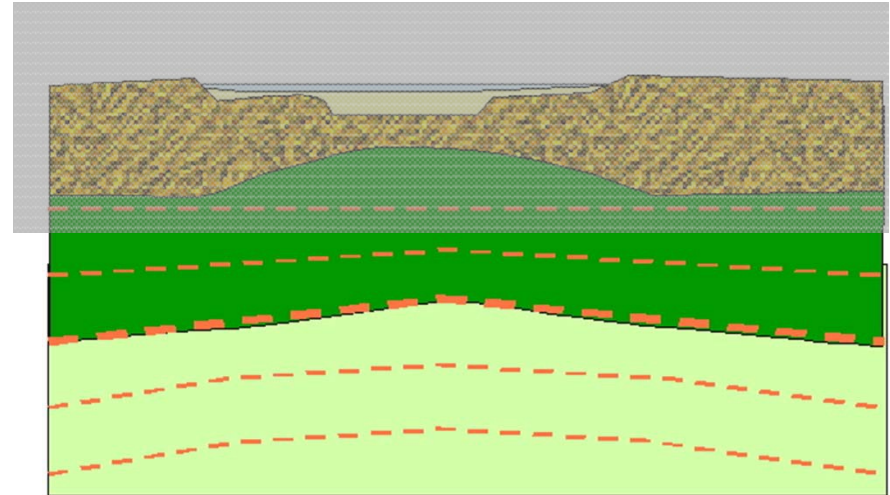
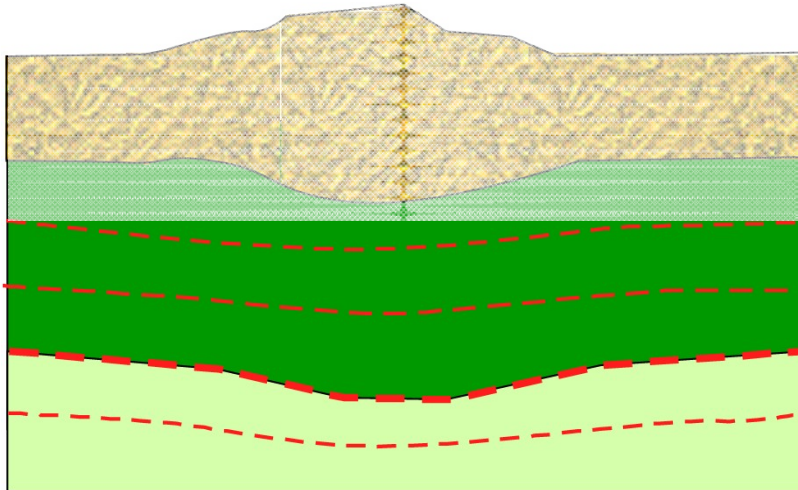
when thinning is **fast**  
 advection  $\gg$  conduction  $\blacktriangleright$   
 strong thermal anomaly  $\blacktriangleright$   
**strong lithospheric thinning**

Back to our cartoons of the **lithospheric mantle**

### Fast deformation

No thickness changes = little/no vertical movements driven by the lithospheric mantle

The extension case

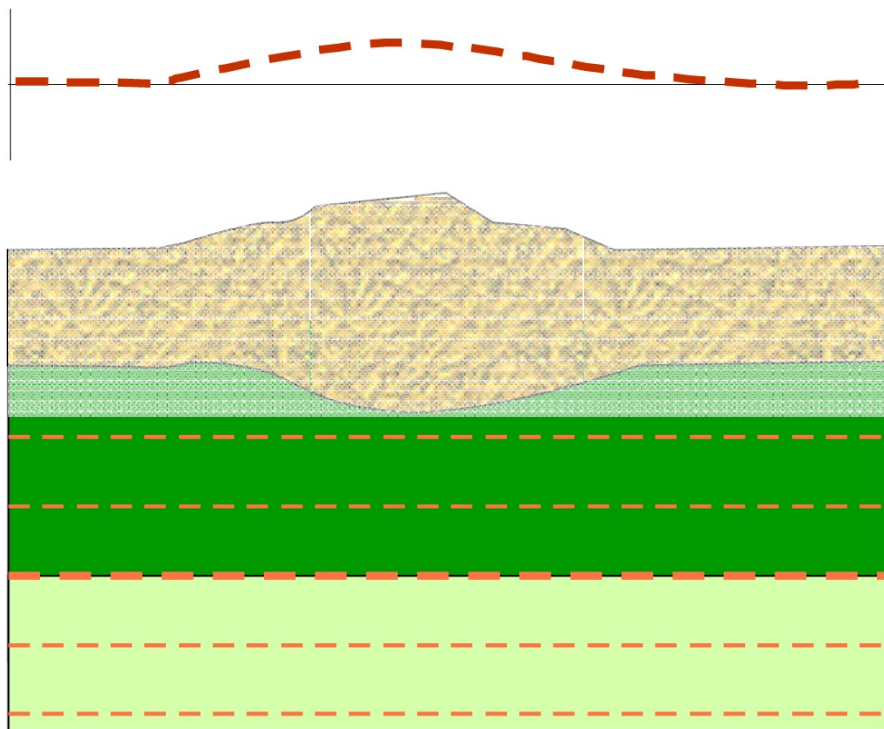


The contraction case

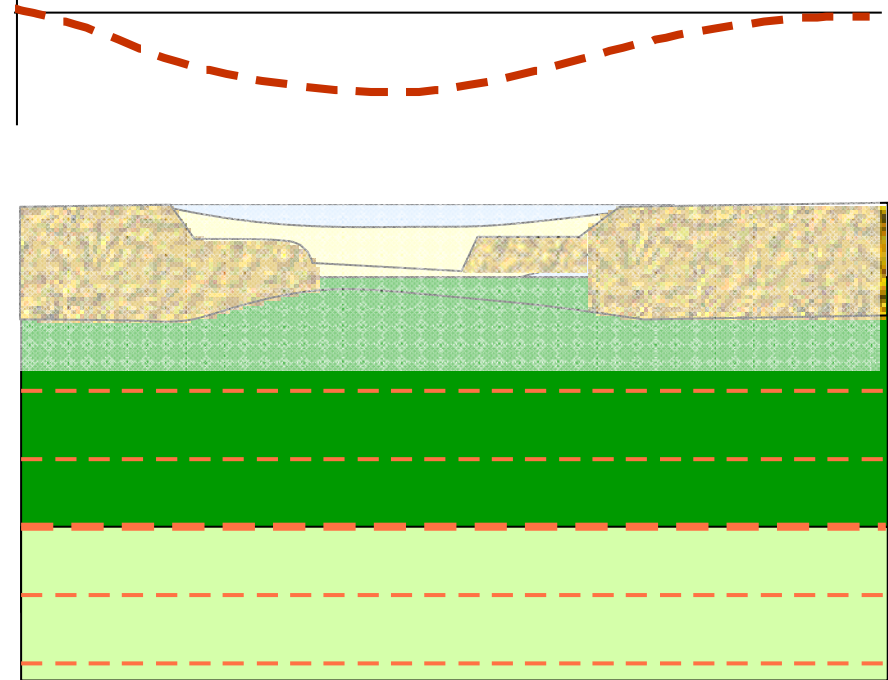


# The lithospheric mantle **during** slow deformation: **Important thickness changes**

In the case of **extension**  
the lithospheric mantle **thickens!**



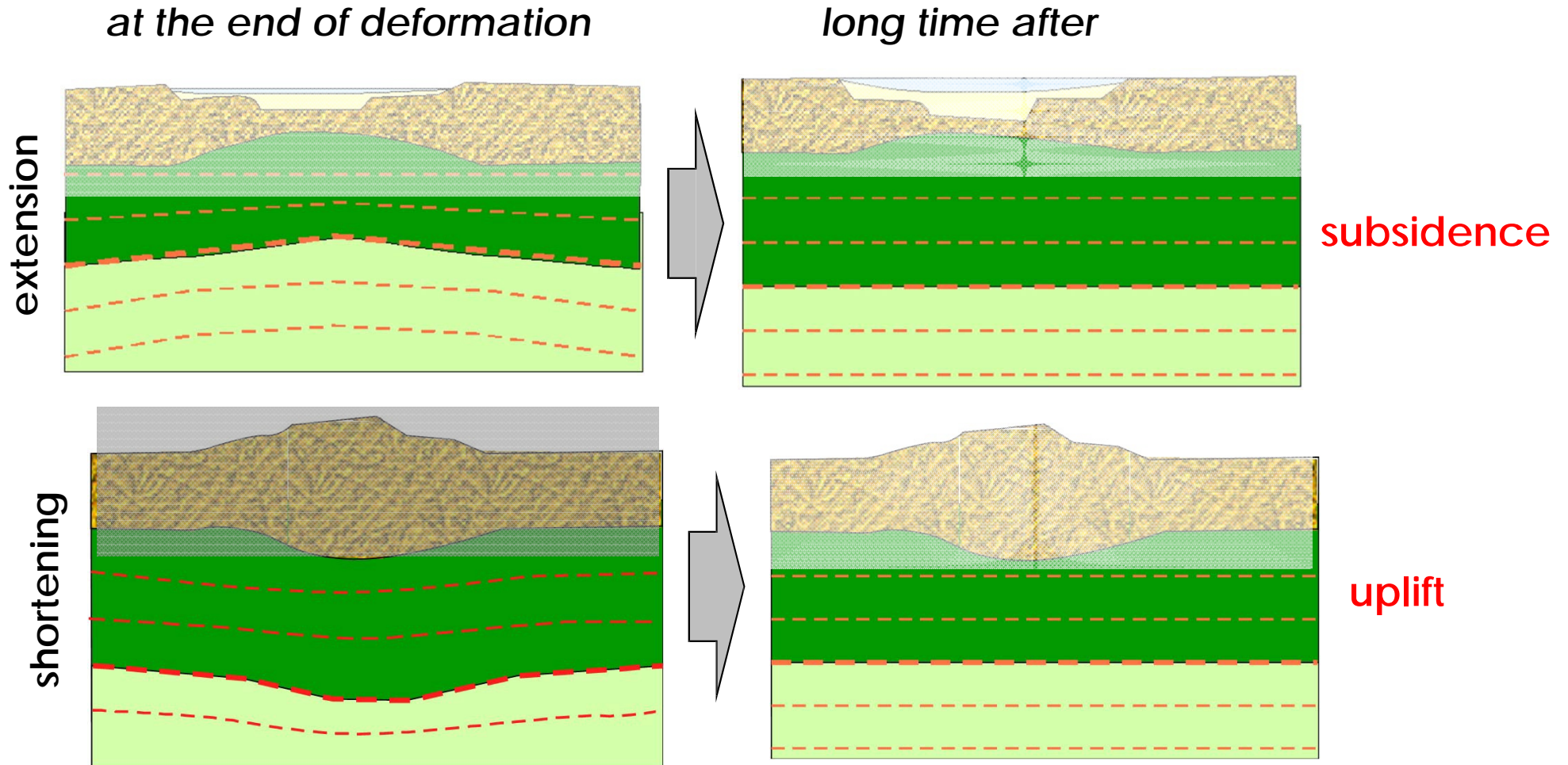
*vertical movements*



In the case of **contraction**  
the lithospheric mantle **thins and**  
**pushed the Earth surface up!**

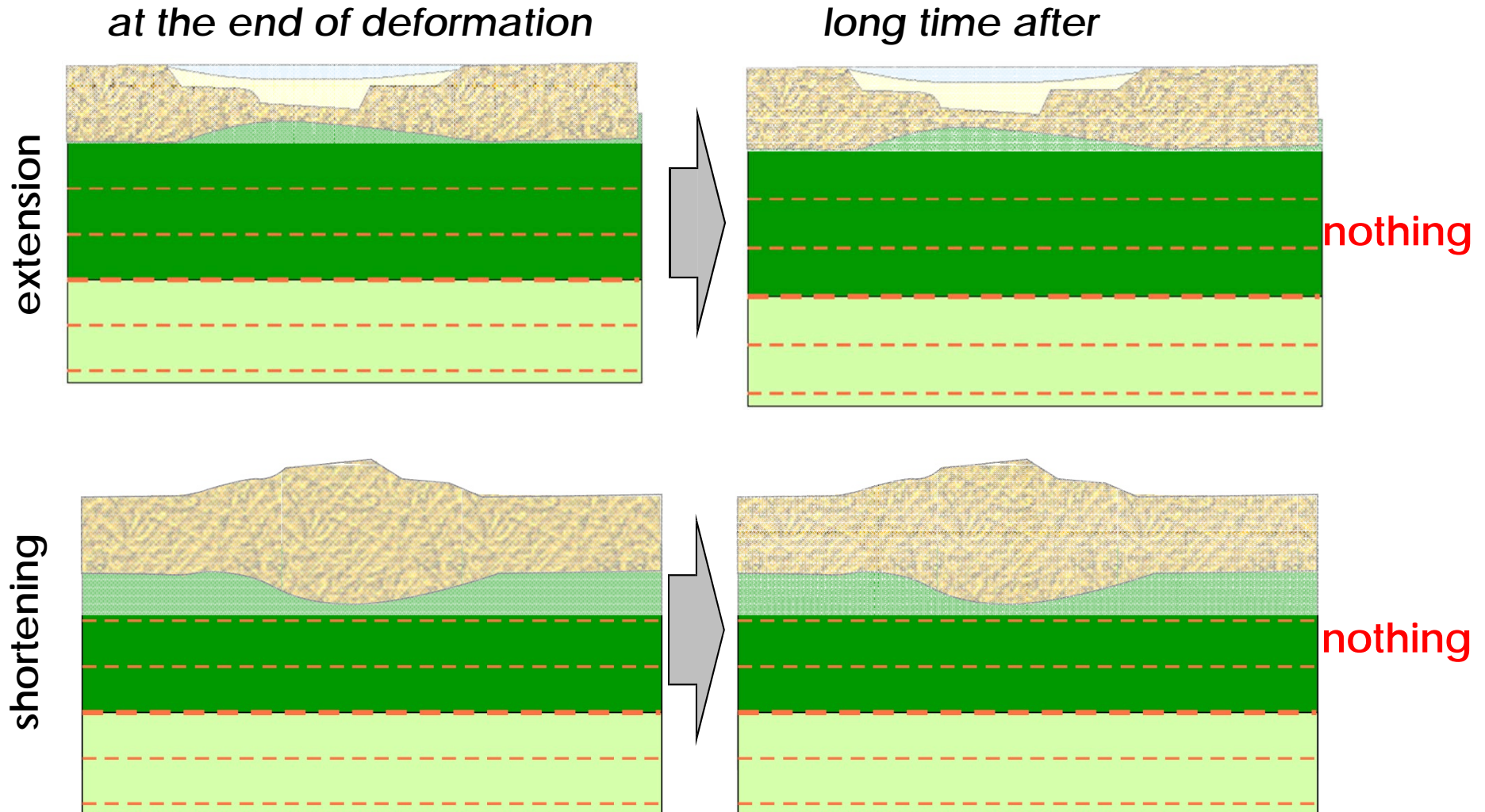
Following the **end** of deformation: **Fast deformation**

The thickness of the lithospheric mantle changes and vertical movements occur



Following the **end** of deformation: **slow deformation**

No changes in the thickness of the lithospheric mantle and no vertical movements occur



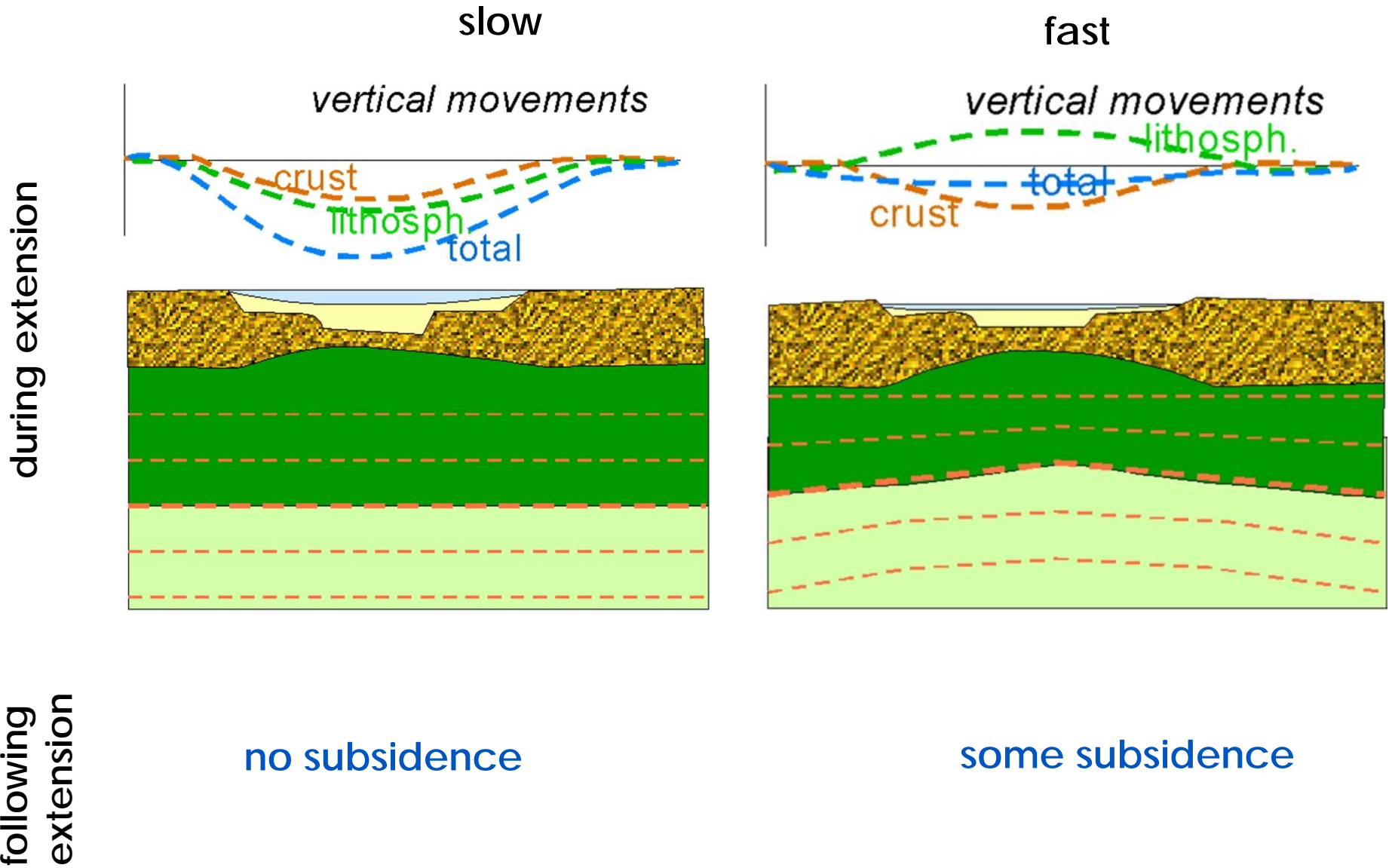
The rates of deformation controls the magnitude of the thermal anomaly present at the end of deformation

The magnitude of vertical movements depends on the amplitude of the thermal anomaly

The faster extension/shortening, the stronger will be vertical movements after the end of deformation

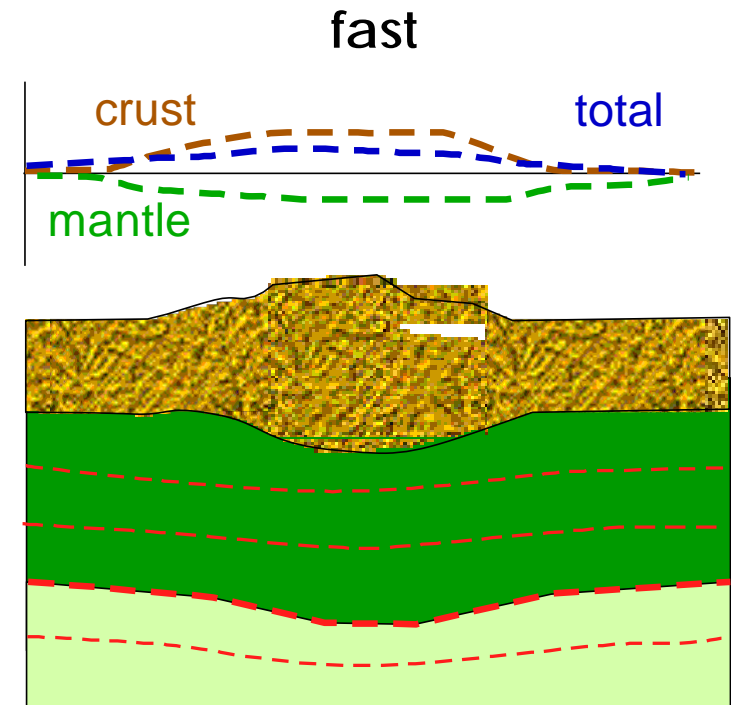
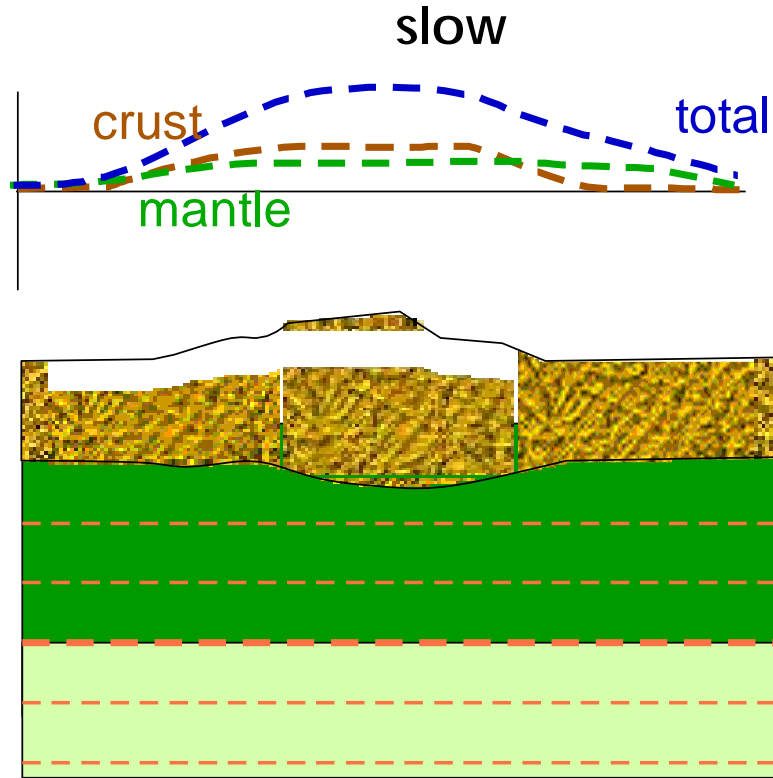


# The real world: combining crust and lithospheric mantle



During contraction/thickening

during contraction



following contraction

no uplift

some uplift

Recap:

during thickness changes:

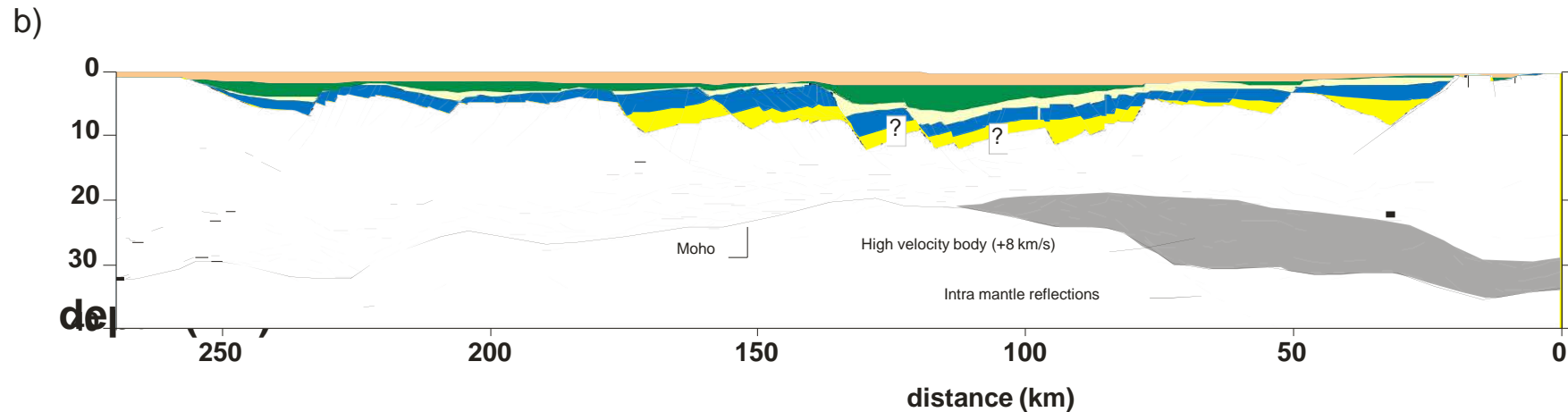
- the crust simply wants subsidence/uplift
- what the lithospheric mantle wants ... depends on the rate of thinning/thickening

Following thickness changes:

- the crust has nothing more to say (neglecting erosion)
- the lithospheric mantle will impose vertical movements; their magnitude depends on the amplitude of the thermal anomaly present at the end of thinning/thickening

What is the meaning of “slow” and “fast”?  
No general answer is possible without knowledge of the geometry of the system and numerical modelling



Processes leading to **load** changesTectonics (**thinning**)

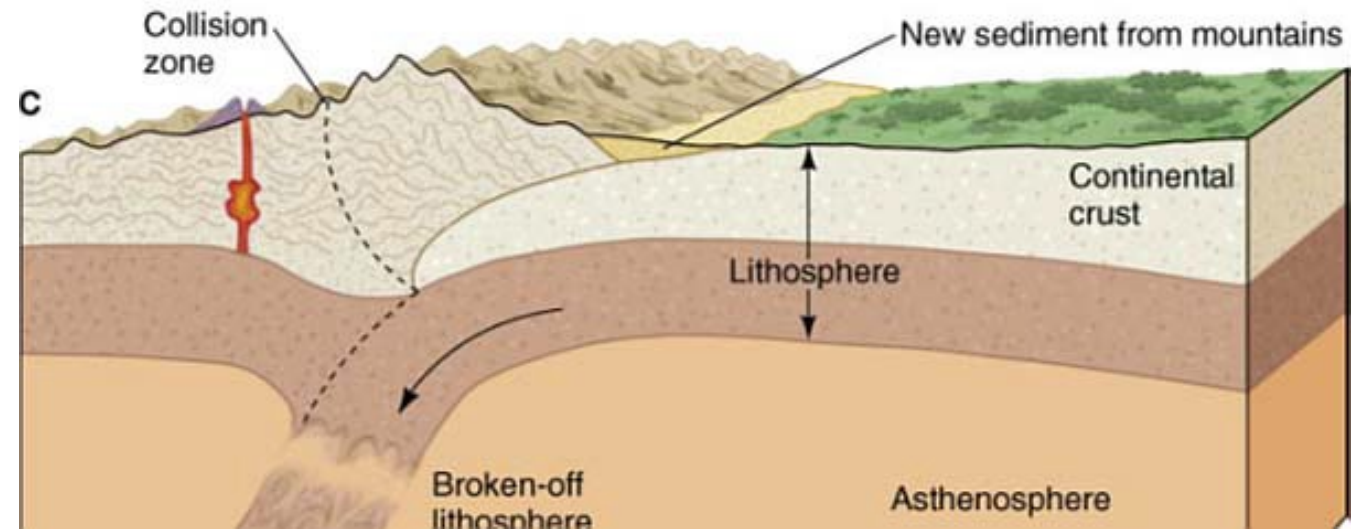
Where is the site of **maximum thinning**?

Where is the site of **maximum subsidence**?

Is there correspondence between the two?

Knowing that the **green sediments** were deposited until 70Ma, what will be the shape of the LAB?

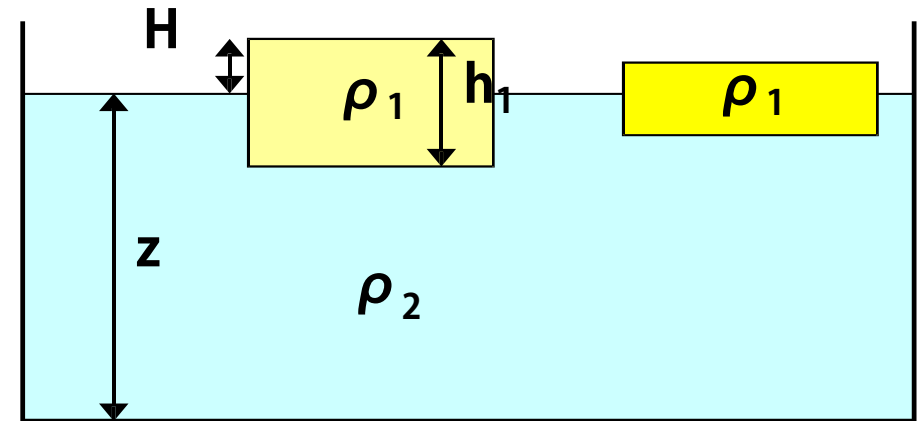


Tectonics (**thickening**)

other **less common** processes can change the **density** of the rocks. One of them is **metamorphism**

# Erosion

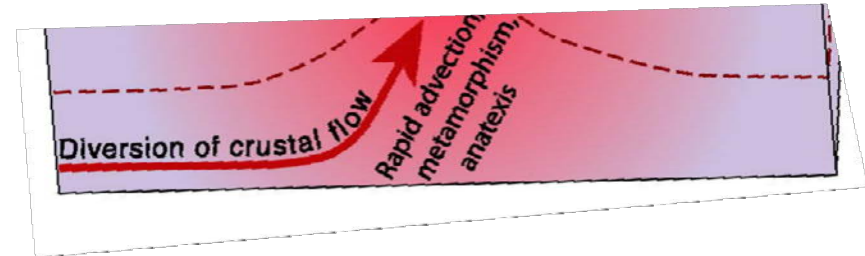
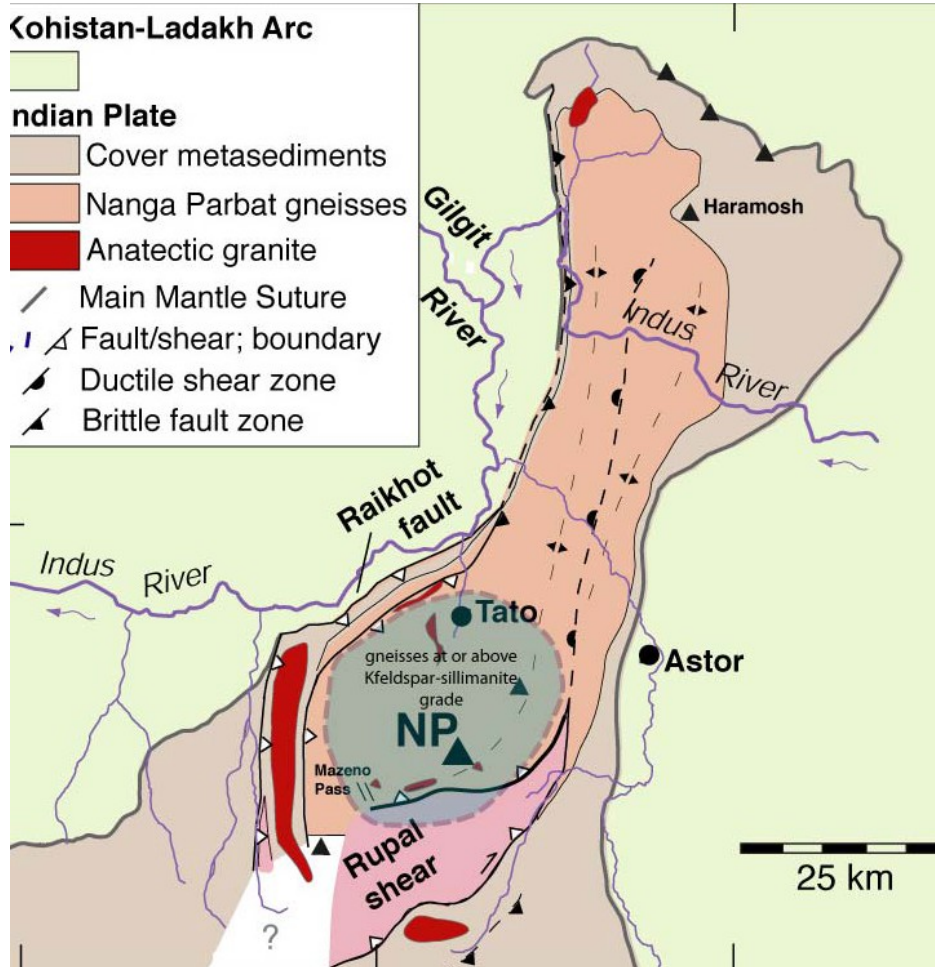
When material is eroded, the surface of the Earth goes **down** but the lower part of the block (the crust) **comes up**!



This leads to less mountain, less erosion and eventually an equilibrium profile

# An extreme situation

Valleys are the places where the strongest erosion takes place



The rivers "pull up" rocks from below

## The Indus river valley

Rocks are found along the river which were formed at depth of >10km and have been brought to the surface in very recent times

# Sources of figures

<http://tsjok45.wordpress.com/2013/06/24/geologie-trefwoord-c/>

[http://bc.outcrop.org/GEOL\\_B10/Dlecture10.html](http://bc.outcrop.org/GEOL_B10/Dlecture10.html)

<http://dc401.4shared.com/doc/Ch0Ecp8M/preview.html>

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[http://www.ees.lehigh.edu/groups/corners/galleries/mapdata/pages/nanga\\_map.html](http://www.ees.lehigh.edu/groups/corners/galleries/mapdata/pages/nanga_map.html)